

RESEARCH LETTER

10.1002/2016GL070710

Key Points:

- Current terrestrial biosphere models underestimate northern CO₂ seasonal cycle amplitude increase
- Models capture observed greening trends and therefore increased uptake of light by ecosystems
- Reconciling these suggests an increase in ecosystem light use efficiency larger than models simulate

Supporting Information:

- Supporting Information S1

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Citation:

Thomas, R. T., et al. (2016), Increased light-use efficiency in northern terrestrial ecosystems indicated by CO₂ and greening observations, *Geophys. Res. Lett.*, 43, 11,339–11,349, doi:10.1002/2016GL070710.

Received 23 JAN 2016

Accepted 12 OCT 2016

Accepted article online 15 OCT 2016

Published online 4 NOV 2016

Increased light-use efficiency in northern terrestrial ecosystems indicated by CO₂ and greening observations

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Abstract Observations show an increasing amplitude in the seasonal cycle of CO₂ (ASC) north of 45°N of 56 ± 9.8% over the last 50 years and an increase in vegetation greenness of 7.5–15% in high northern latitudes since the 1980s. However, the causes of these changes remain uncertain. Historical simulations from terrestrial biosphere models in the Multiscale Synthesis and Terrestrial Model Intercomparison Project are compared to the ASC and greenness observations, using the TM3 atmospheric transport model to translate surface fluxes into CO₂ concentrations. We find that the modeled change in ASC is too small but the mean greening trend is generally captured. Modeled increases in greenness are primarily driven by warming, whereas ASC changes are primarily driven by increasing CO₂. We suggest that increases in ecosystem-scale light use efficiency (LUE) have contributed to the observed ASC increase but are underestimated by current models. We highlight potential mechanisms that could increase modeled LUE.

1. Introduction

Observations show that the terrestrial biosphere is responding to anthropogenic environmental change [Friedlingstein et al., 2014; Ciais et al., 2014]. Photosynthetic CO₂ uptake currently exceeds release through respiration and other processes, and the terrestrial biosphere removes around a quarter of anthropogenic CO₂ emissions each year [Ciais et al., 2014]. Net uptake of CO₂ can be partly attributed to “CO₂ fertilization,” since increases in ambient CO₂ can increase the biochemical rate of photosynthesis and increase water use efficiency [Franks et al., 2013; Penuelas et al., 2011; Keenan et al., 2013]. Net uptake of CO₂ may also result in part from changes in climate, land management, ecosystem composition, and N deposition, among other factors [Huntzinger et al., 2013]. Better understanding of how the terrestrial biosphere has responded to changes in climate, CO₂, and other drivers over recent years is needed to understand how it might change in the future and whether it will continue to serve as a net sink [Le Quere et al., 2015; Friedlingstein et al., 2006; Schimel et al., 2015]. In this paper we compare long-term observations of atmospheric CO₂ and greenness to terrestrial biosphere model output.

The seasonal cycle of CO₂ in the Northern Hemisphere can be almost entirely attributed to terrestrial biosphere exchange [Randerson *et al.*, 1997; Graven *et al.*, 2013], so changes in the seasonal cycle of CO₂ indicate large-scale changes in northern terrestrial ecosystem exchange. Increases in the amplitude of the seasonal cycle of CO₂ (ASC) have been observed at the measurement stations Mauna Loa, Hawaii (Mauna Loa Observatory (MLO)), and Point Barrow, Alaska (BRW) of $0.32 \pm 0.07\% \text{ yr}^{-1}$ and $0.60 \pm 0.05\% \text{ yr}^{-1}$, respectively, since the early 1960s [Keeling *et al.*, 1996; Graven *et al.*, 2013]. Recently, Graven *et al.* [2013] measured long-term changes in ASC throughout the higher latitudes of the Northern Hemisphere. Comparisons between two aircraft campaigns in 1958–1961 and 2009–2011 showed an increase in ASC of $5.0 \pm 0.8 \text{ ppm}$ ($56 \pm 9.8\%$) between 45 and 90°N at 500 mb, indicating that the seasonal uptake and release of CO₂ in terrestrial ecosystems between 30 and 90°N increased by 32–59% over this time [Graven *et al.*, 2013]. The largest contribution to the ASC was in the main growing season [Graven *et al.*, 2013]. Increases in ASC may therefore be related to the extratropical net carbon sink as studies suggest that increases in net terrestrial carbon uptake during the growing season drive the northern extratropical net carbon sink [Rayner *et al.*, 2015; Gurney and Eckels, 2011].

Vegetation greenness, defined here as the fraction of photosynthetically active radiation that is absorbed by green vegetation on land (fAPAR), has been measured by satellites since mid-1981 [Murray-Tortarolo *et al.*, 2013]. An increase of vegetation greenness has been observed between 30 and 90°N of 8.7% for all types of vegetation from 1982 to 2010 (Figures 4b and S3b in the supporting information). These trends imply that photosynthetic activity in northern latitudes has increased over the last three decades [Slayback *et al.*, 2003; Zhu *et al.*, 2013]. Many studies attribute greening primarily to rising temperature in spring and autumn, leading to a reduction in snow cover, an earlier start to the growing season [Piao *et al.*, 2008; Myneni *et al.*, 1997; Notaro *et al.*, 2005; Xu *et al.*, 2013] and increased peak greenness [Buitenwerf *et al.*, 2015; Myneni *et al.*, 1997].

Observations of the CO₂ seasonal cycle and vegetation greenness, and their changes over time, present opportunities to evaluate terrestrial biosphere models. Reproducing the mean CO₂ seasonal cycle requires that modeled photosynthesis and respiration have the correct phase and magnitude. Reproducing the changes in ASC additionally requires that interactions between the terrestrial biosphere and the environmental drivers, including climate, CO₂, land use, and disturbance (e.g., fire and insects), are well represented in models [McGuire *et al.*, 2001]. Previous comparisons have shown that the long-term increase in ASC observed in the aircraft data was underestimated by the coupled carbon-climate CMIP5 Earth System Models [Graven *et al.*, 2013], whereas the smaller trend observed in ASC at ground-based NOAA stations over 1982–2011 was captured by one off-line dynamic vegetation model (LPJml). In LPJml, the ASC increase was primarily driven by increases in vegetation cover and plant productivity resulting from increasing temperatures at high latitudes [Forkel *et al.*, 2016]. However, other studies that examined seasonal biosphere fluxes have found conflicting results. For the off-line terrestrial biosphere models taking part in the Multiscale Synthesis Terrestrial Model Intercomparison Project (MsTMIP) and the TRENDY dynamic vegetation model intercomparison project, changes in seasonal fluxes were predominately driven by climate change in only a minority of models. In most current biosphere models, the increase in atmospheric CO₂ is the primary driver of changes in seasonal fluxes [Ito *et al.*, 2016; Zhao *et al.*, 2016]. In comparisons of observed and modeled greening trends, the TRENDY models suggest that greening has been primarily driven by climate change at higher northern latitudes (60–90°N) and by CO₂ at lower northern latitudes (30–60°N) [Zhu *et al.*, 2016].

There is a clear need to identify mechanisms that reconcile observed and modeled changes in ASC and greening and to clarify the role of CO₂, climate change, and other factors driving ecosystem changes. A key distinction can be made between changes in the structure (leaf area) and the physiology (intrinsic rate of carbon uptake) of vegetation by using the concept of light use efficiency (LUE), the net primary productivity (NPP) per unit of light absorbed (APAR) [Norby *et al.*, 2003]. If increases in NPP for a given ecosystem were driven solely by increased leaf area, and therefore increased APAR, NPP would relate to APAR by a constant factor corresponding to the LUE of the ecosystem. However, increases in NPP may also arise from increases in LUE, even in the absence of APAR increases. A decade of plot-scale observations from four temperate sites in the Free Air CO₂ Enrichment experiments show that in medium- to high-density canopies, increases in NPP were driven by LUE increases, while there was no significant change in APAR [Norby *et al.*, 2003; McCarthy *et al.*, 2006; Long *et al.*, 2004]. Analysis of LUE trends in models can therefore help to elucidate mechanisms of ecosystem change and to inform model-data comparisons of fAPAR and CO₂ exchange [Norby *et al.*, 2005].

In this paper, we compare observed changes in ASC and greening trends to output from 13 biosphere models in MsTMIP, applying the concept of LUE to interpret the model-data comparison. In MsTMIP, the process-based

terrestrial biosphere models ran semifactorial experiments where photosynthesis and respiration were driven by a standard set of observed changes in climate, land use, CO₂, and in some cases N deposition [Huntzinger *et al.*, 2013; Wei *et al.*, 2014]. We compare modeled output of greenness to the longest available record of satellite fAPAR data—fPAR3g [Zhu *et al.*, 2013]. We use the TM3 atmospheric transport model with net ecosystem production from MsTMIP to evaluate the modeled seasonal cycle amplitude of atmospheric CO₂ directly against aircraft measurements in 1958–1961 and 2009–2010. The aircraft data are more representative of large-scale behavior than surface measurement stations [Sweeney *et al.*, 2015; Keppel-Aleks *et al.*, 2012; Sawa *et al.*, 2012], showing a clearer and more consistent signal. Thus, aircraft data are a stronger constraint on large-scale ecosystem changes than the long-term records at the surface stations MLO and BRW, but they have previously been compared only with CMIP5 models [Graven *et al.*, 2013].

We show that most MsTMIP models simulate overall greening trends well but underestimate the observed ASC change. Increases in modeled NPP in northern ecosystems are a combination of CO₂-driven increases in LUE and climate-driven greening. Since models capture greening trends but underestimate the change in ASC, we suggest that models underestimate increases in ecosystem LUE over the past several decades and we discuss potential causes for this phenomenon.

2. Methodology

2.1. Observed Data

2.1.1. Seasonal Amplitude of Atmospheric CO₂

The aircraft campaigns “International Geophysical Year” in 1958–1961 and “HIAPER Pole-to-Pole Observations” in 2009–2011 observed atmospheric CO₂ in the middle to lower troposphere [Graven *et al.*, 2013]. We focus on the change in the amplitude of the seasonal cycle of CO₂ (ASC) at 500 mb in the 45–90°N region, where the largest increase in ASC was observed: 5.0 ± 0.8 ppm ($56 \pm 9.8\%$). We do not consider phase shifts in the CO₂ seasonal cycle, since they are not clearly resolved either by aircraft observations [Graven *et al.*, 2013] or monthly model output.

2.1.2. Vegetation Greenness

Vegetation greenness was analyzed using the remotely sensed third-generation fAPAR data product, fPAR3g, at 1/12° resolution [Zhu *et al.*, 2013]. The data were averaged to the 0.5° × 0.5° grid of the MsTMIP models over 1982–2011 [Zhu *et al.*, 2013]. Monthly fAPAR was calculated as the mean of 15 day cloud-free composites, then weighted by average monthly PAR per grid cell and summed for months with a temperature above 0°C to get annual growing season fAPAR (GS-fAPAR). Total GS-fAPAR was calculated as the mean of all land grid cells north of 30°N. See Text S1 for full methods.

2.2. Model Output

2.2.1. MsTMIP Models

We analyze CO₂ fluxes from 13 terrestrial biosphere models in the MsTMIP V1.0 experiments (Table S1) [Huntzinger *et al.*, 2014]. Global monthly model output is at 0.5° × 0.5° resolution. Analysis was carried out on years overlapping with the aircraft and satellite observation periods: 1958–1961 and 2009–2010, and 1982–2010 respectively. A standard protocol and environmental driver data set were used to spin-up models to steady state, and control simulations were run for each model with environmental drivers held at preindustrial levels. Experiments were run in which historical climate, land use change (LUC), and CO₂ were sequentially included as model drivers and then N deposition for BIOME-BGC, CLM4, CLM4VIC, DLEM, ISAM, and TEM6 [Huntzinger *et al.*, 2013]. The historical drivers consist of observed and reanalysis data. A land cover map was reinterpreted by each modeling team, for use as the LUC driver [Wei *et al.*, 2014]. Since the models are driven by the same data, intermodel differences reflect how terrestrial biosphere processes have been represented in each model [Dalmonech and Zaehle, 2013; Le Quere *et al.*, 2015; Huntzinger *et al.*, 2013; Sitch *et al.*, 2015]. Unless otherwise stated, the simulation including all time-varying historical drivers is used in the results. We also analyze individual contributions to fluxes from climate, LUC, CO₂ or N deposition, which were calculated as the difference between the first simulation in which that type of historical data was used and the previous simulation. One model, BIOME-BGC, only ran the “climate-only” and “all-on” experiments, with fixed LUC and N deposition.

Monthly net ecosystem productivity (NEP) was calculated from model outputs as the difference between net primary production and heterotrophic respiration (NPP-Rh) [Chapin *et al.*, 2006]. This neglects fire and other disturbances that are often absent or poorly represented in models and poorly constrained by observations [Huntzinger *et al.*, 2013]. The CO₂ concentration resulting from the NEP fluxes was calculated for each

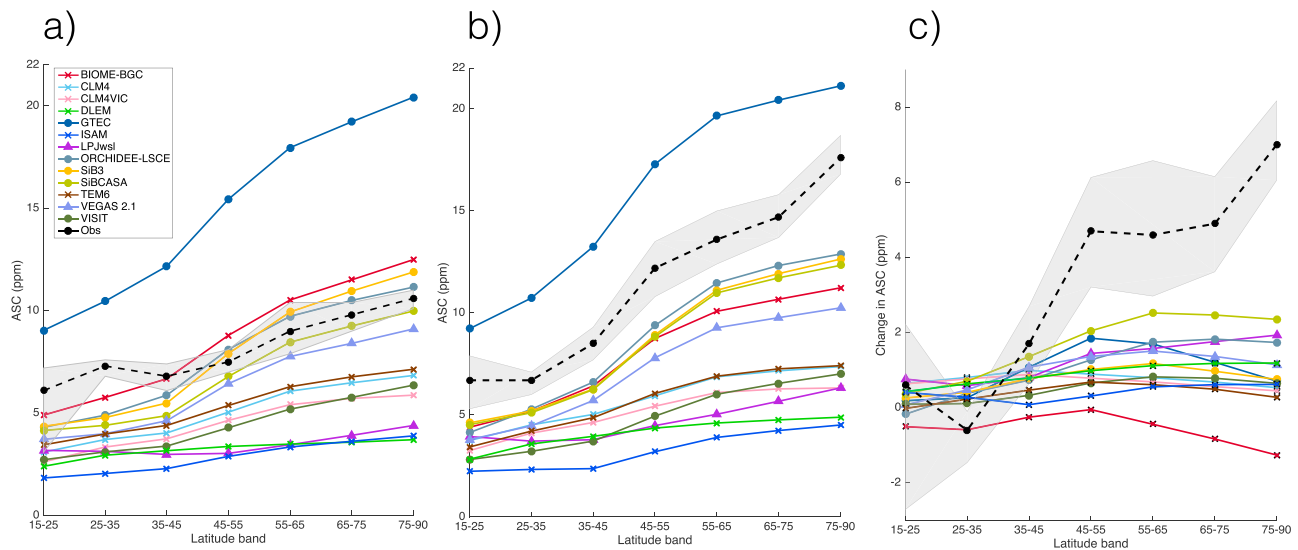


Figure 1. Observed and modeled ASC by latitude for (a) 1958-61, (b) 2009-10 and (c) difference between 2009-10 and 1958-61. Model type depicted by markers: no nitrogen cycle, no dynamic vegetation (circles), no nitrogen cycle, dynamic vegetation (triangles), and nitrogen cycle, no dynamic vegetation (crosses). Grey shading is observational uncertainty from [Graven *et al.*, 2013].

model and simulation using the TM3 atmospheric transport model at $5^\circ \times 3.83^\circ$ resolution, forced by National Centers for Environmental Prediction reanalysis data specific for each year [Heimann and Korner, 2003]. This captures any changes in atmospheric transport, even though transport effects were shown to have little impact on long-term ASC trends [Graven *et al.*, 2013]. Compared to other atmospheric transport models, TM3 does not show large biases in seasonal CO_2 amplitude or vertical exchange [Gurney *et al.*, 2004; Stephens *et al.*, 2007]. The simulated CO_2 concentration from NEP was added to simulated CO_2 concentrations from monthly ocean [Patra *et al.*, 2011; Roedenbeck *et al.*, 2003] and fossil fuel [Andres *et al.*, 2011] fluxes for the same years. The total CO_2 concentration for each model was then detrended for each latitude band using a polynomial fit, and the mean CO_2 concentration was subtracted and interpolated to aircraft data at 500 mb.

Monthly modeled fluxes of NPP, Rh, and NEP were each summed over $30\text{--}90^\circ\text{N}$ for 1958–1961 and 2009–2010, and the seasonal amplitude and mean annual flux was calculated for each flux and time period. The flux and CO_2 concentration amplitudes were calculated as the difference between the maximum and minimum of the mean seasonal cycle.

Modeled leaf area index (LAI, the projected area of leaves over an area of ground [Murray-Tortarolo *et al.*, 2013]) was converted to fAPAR using the Beer’s Law approximation:

$$\text{fAPAR} = 1 - e^{-0.5\text{LAI}} \quad (1)$$

Modeled GS-fAPAR was calculated in the same way as observed GS-fAPAR (section 2.1.2). Modeled growing season LUE was calculated for land $30\text{--}90^\circ\text{N}$ in 1958–1961 and 2009–2010 as

$$\text{GS-LUE} = \frac{\text{GS-NPP}}{\text{GS-APAR}}, \quad (2)$$

where

$$\text{GS-APAR} = \text{GS-fAPAR} \times \text{PAR}. \quad (3)$$

3. Results

3.1. Seasonal CO_2 Amplitude Change

MtMIP models are generally able to reproduce the amplitude of the CO_2 seasonal cycle and its latitudinal gradient at 500 mb in 1958–1961 but not in 2009–2010 (Figure 1). Models simulate a wide range of ASC change north of 45°N , -0.5 to 2.4 ppm, all substantially lower than observed (Figures 1 and S2). The models generally

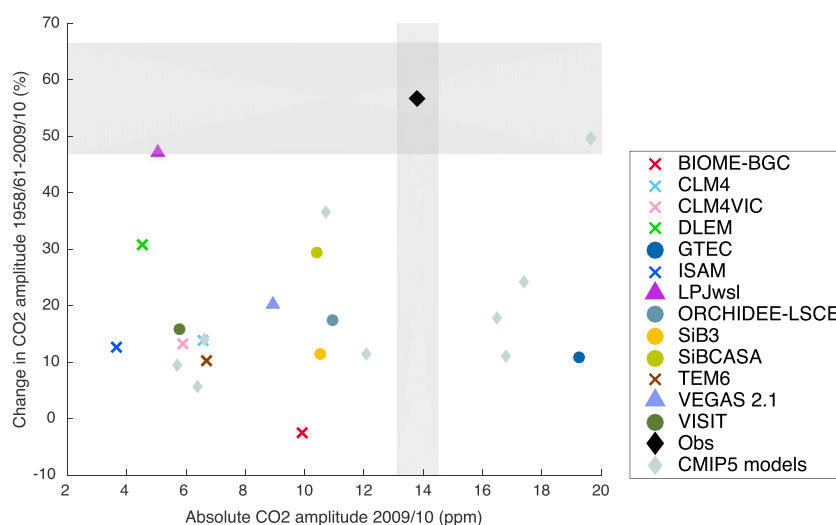


Figure 2. ASC in 2009–2010 compared to percentage change in ASC from 1958–1961 to 2009–2010 for aircraft observations (black square), MsTMIP models (colored markers), and CMIP5 models for TM3 only (grey markers) (revised from [Graven *et al.*, 2013]). Grey shading is observational uncertainty [Graven *et al.*, 2013].

capture the phase of the CO₂ seasonal cycle (Figure S3a). The primary driver of the increase in modeled ASC is CO₂ (Figure 4). Coupled carbon-nitrogen cycle (C-N) models have among the smallest changes in ASC and the weakest CO₂-driven increases. The small CO₂-driven ASC increase in C-N models is likely due to N limitation, as ASC further increases in simulations when time-varying N deposition is included. LPJwsl is the only model that has a fractional increase in ASC within the observational uncertainty (47%) (Figure 2), but its absolute ASC and ASC change are much too small (Figure 1). The ASC change in the MsTMIP models is similar to the CMIP5 models [Graven *et al.*, 2013, Figure 2], although MsTMIP models tend to have smaller 2009–2010 amplitudes.

3.2. Vegetation Greenness

Models are generally able to reproduce the observed interannual variability in GS-fAPAR and its increase of 0.02 (8.7%) between 30 and 90°N from 1982 to 2010 (Figures 4b and S3b). Observed and modeled greening was due to an increase in leaves per unit area, rather than vegetated area. Only ISAM showed an increase in vegetated area. All models show that greening is driven by climate (Figure 4b), and models perform best in arctic and boreal regions where greening trends have been driven by increasing temperatures [Piao *et al.*, 2006] (Figure S1). There is little influence on the overall greening trend from LUC, CO₂, and N deposition. These drivers may be of more importance at lower latitudes where models do not match observed trends well (Figure S1) and where greening trends depend more on how vegetation is represented in models [Eastman *et al.*, 2013; Murray-Tortarolo *et al.*, 2013; Zhu *et al.*, 2016].

3.3. Terrestrial Carbon Fluxes

3.3.1. NEP

The amplitude of NEP varies from less than 1 Pg C k yr⁻¹ to more than 5 Pg C yr⁻¹, consistent with previous studies showing MsTMIP models have high intermodel variability in NEP, as well as in its constituent fluxes, gross primary production, and ecosystem respiration [Huntzinger *et al.*, 2013; Zscheischler *et al.*, 2014; Schwalm *et al.*, 2015]. NEP amplitude generally increased (Figure 3b), as shown previously for MsTMIP [Ito *et al.*, 2016] and TRENDY [Zhao *et al.*, 2016]. NEP amplitude is well correlated to the simulated ASC across the models; $r^2=0.95$ (Figure 3a). There is also a good correlation between NEP amplitude change and ASC change (Figure 3b), though somewhat weaker ($r^2=0.70$), which may be due to additional influences on ASC change from fluxes south of 30°N.

Figure 3b also shows the optimized flux of NEP, from Graven *et al.* [2013], where NEP amplitude in arctic, boreal, and temperate regions were adjusted to best match the pattern of observed ASC change in aircraft and surface data. The optimized increase in NEP amplitude from Graven *et al.* [2013] is larger than any modeled increase, but it lies within the uncertainty of the model regression line, indicating that the relationship between ASC and NEP amplitude change is similar.

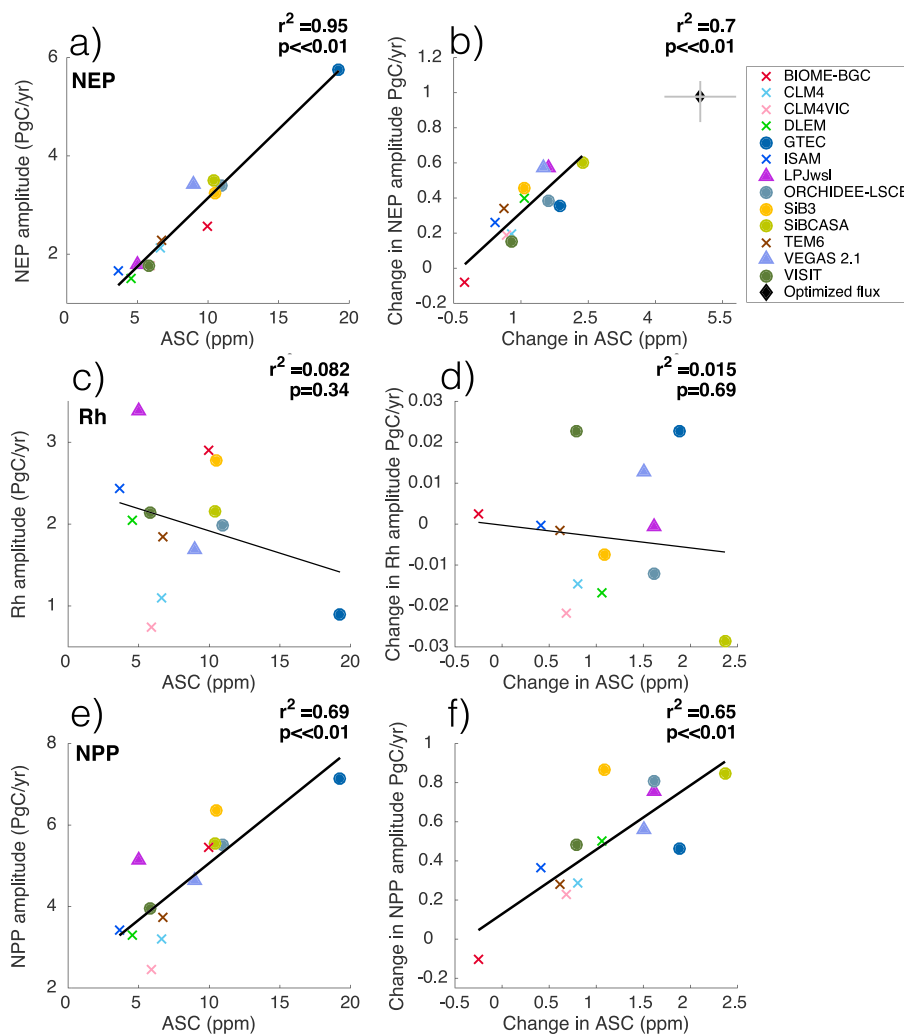


Figure 3. Modeled ASC and amplitude of (a) NEP, (c) NPP and (e) Rh for 2009–2010. This relationship is similar in 1959–1961. Modeled ASC change and change in amplitude of (b) NEP, (d) NPP, and (f) Rh. All panels except Figure 3b show fluxes for land 30°N–90°N and ASC for 45°N–90°N at 500 mb and regression lines are plotted, with markers as in Figure 1. In Figure 3b, NEP is for arctic, boreal, and temperate regions only, defined by the biome mask in *Graven et al.* [2013]. The black diamond is observed ASC change and optimized flux of NEP from *Graven et al.* [2013].

3.3.2. Rh

There is no significant correlation between modeled Rh amplitude and ASC or between Rh amplitude change and ASC change ($r^2 \approx 0$) (Figures 3e and 3f); thus, changes in Rh cannot be driving the change in modeled NEP amplitude and ASC. Mean annual Rh increased, but changes in Rh amplitude were small because Rh increased throughout the year rather than in the month of maximum Rh (Figures 3f and 4c). The small changes in MsTMIP Rh amplitude are inconsistent with CMIP5 models, which show much larger Rh amplitude changes [*Graven et al.*, 2013, Figure S9], perhaps due to climate feedbacks in CMIP5 models that are not present in the off-line MsTMIP models.

The largest increase in mean annual Rh is generally from CO₂, although climate is dominant in VEGAS, LPJwsl, and N deposition is dominant in CLM4VIC and CLM4 (Figure 4c).

3.3.3. NPP

The modeled seasonal amplitude of NPP and change in NPP amplitude are significantly correlated to modeled ASC and change in ASC across the models; $r^2 = 0.69$ and $r^2 = 0.65$, respectively (Figures 3c and 3d). Mean annual NPP and seasonal NPP amplitude increased for all models between the two periods (except NPP amplitude in BIOME-BGC) (Figure 4c). Hence, changes in modeled NPP, not Rh, are driving the change in modeled NEP amplitude and ASC.

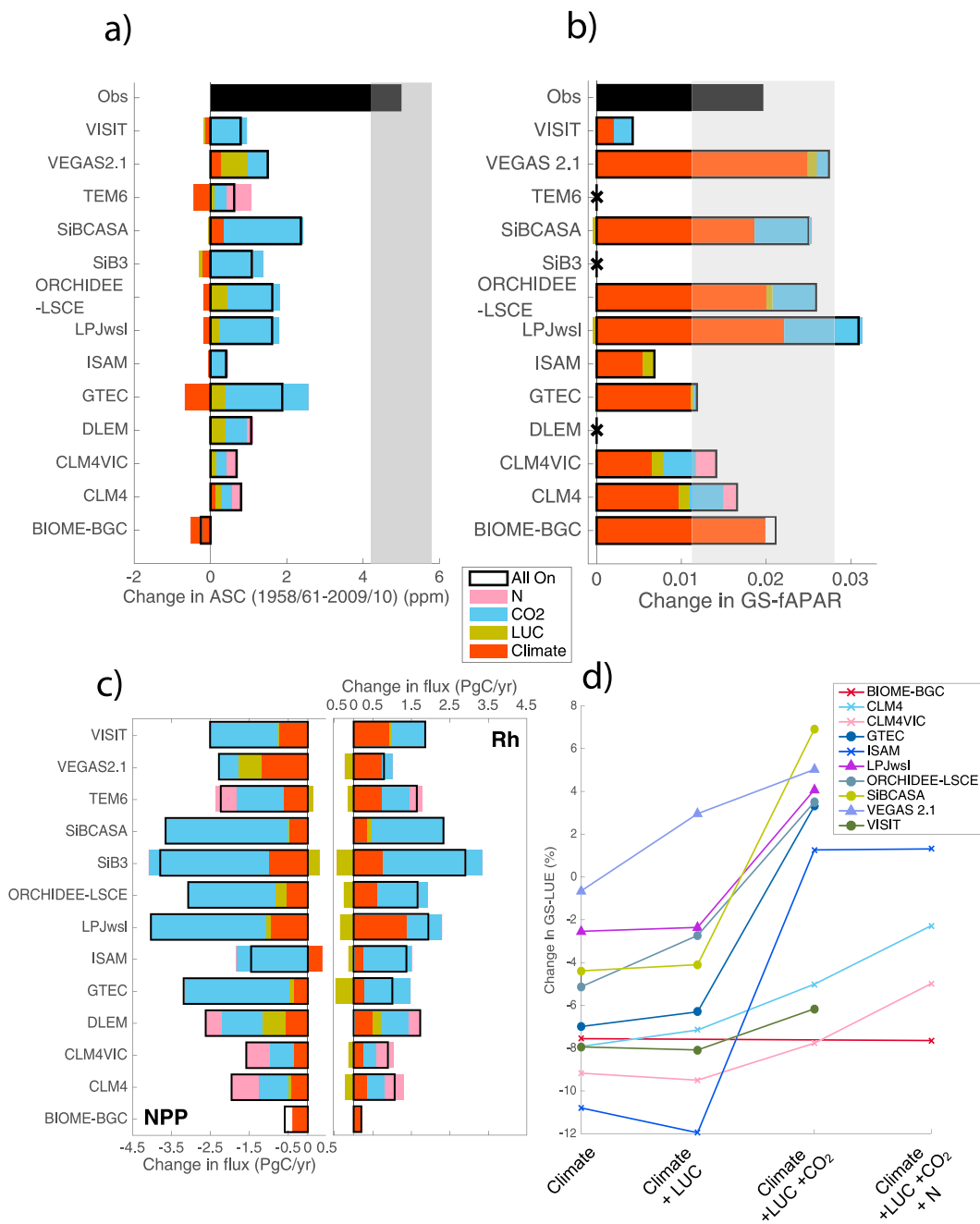


Figure 4. Contributions from climate, LUC, and CO₂ N deposition to modeled changes in (a) ASC, (b) GS-fAPAR, (c) mean annual NPP and Rh, and (d) GS-LUE. ASC is for 45°N–90°N at 500 mb between 1958–1961 and 2009–2010 (Figure 4a). Fluxes are for 30°N–90°N between 1982–2010 (Figure 4b) and 1958–1961 to 2009–2010 (Figures 4 and 4d). Fluxes are positive into the atmosphere. In Figure 4a, observations and uncertainty (grey shading) are from [Graven *et al.*, 2013]. In Figure 4b, the 95% confidence interval in the observed trend is shown (grey shading). LAI was not available for DLEM and TEM6, and LAI was assimilated from remote sensing for SiB3, so no output is shown in Figures 4b or 4d for these models.

CO₂ is the primary driver of NPP increase for most models, while climate and LUC generally result in smaller changes in NPP (Figure 4c). Changes in NPP and NPP amplitude are smallest in C-N models, resulting in small NEP and ASC change (Figures 3b and 4a).

3.4. Light Use Efficiency

Models show mixed results for the overall change in GS-LUE (–7.8 to 6.9%, Figure 4d), but the sign of the LUE response to climate change is consistent across models. Climate change decreases GS-LUE in all models, while

increasing CO₂ increases GS-LUE in all models. Increases in LUE result from N deposition in CLM4 and CLM4VIC, and from LUC in VEGAS2.1 and ORCHIDEE-LSCE, while LUC has little influence on LUE in the other models.

Figure 4 shows that CO₂ is driving increases in modeled NPP, which results in increases in GS-LUE and ASC. On the other hand, while NPP and greening are also increasing with climate change, this does not translate to increased GS-LUE or ASC.

4. Discussion and Conclusion

Current terrestrial biosphere models simulate greening trends reasonably well, but the large increase in observed ASC at high latitudes remains unexplained. There is little difference in simulated ASC or ASC change between models driven with prescribed climate and CO₂ [Zhao *et al.*, 2016; this study], or with modeled climate as in CMIP5 [Graven *et al.*, 2013] (Figure 2), indicating that the disagreement between observations and models lies in the modeling of terrestrial biosphere processes rather than in the climatic forcing. Consistent with previous studies [Graven *et al.*, 2013; Randerson *et al.*, 1997], Figure 3 implies that modeled increases in NEP amplitude lead to increases in modeled ASC. Furthermore, our model-based results suggest NPP as the main driver of NEP amplitude change, also consistent with previous studies [Randerson *et al.*, 1997; Forkel *et al.*, 2016; Zhao *et al.*, 2016]. NEP changes may also be influenced by respiration or disturbance [Zimov *et al.*, 1999], but these effects must be secondary to NPP [Houghton, 1987; Randerson *et al.*, 1997; Graven *et al.*, 2013; Forkel *et al.*, 2016]. Since the MsTMIP models underestimate ASC changes, they are likely also to have underestimated increases in the seasonal amplitude of NPP.

Since NPP can increase through vegetation greening (increased APAR) and/or through an increase in ecosystem LUE (equation (2)), the models' ability to capture the greening trends, and the absence of a trend in PAR (Figure S5), suggests that discrepancies with observed ASC result primarily from errors in the models' representation of LUE. In particular, we suggest that models are capturing the increase in NPP in the early growing season through climate-driven greening but underestimating the GS-LUE-driven increase in NPP in the peak of the growing season when canopies are denser. Our model-data comparison thus provides strong evidence for an increase in GS-LUE across ecosystem in high northern latitudes that is larger than simulated by MsTMIP models.

Increased LUE relates to physiological changes that enhance NPP in addition to, or in the absence of, structural changes through increased leaf area. NPP can increase through CO₂ fertilization, as elevated CO₂ increases photosynthesis and water use efficiency [Ainsworth and Long, 2005; Norby *et al.*, 2003]. CO₂ fertilization is included in the MsTMIP models [Huntzinger *et al.*, 2013], and it is the primary driver of simulated NPP increases (Figure 4c). In models, CO₂ fertilization plays a key role during the peak growing season when canopies are more dense, as seen in Figure 4 in which CO₂-driven NPP increases correspond to GS-LUE and ASC increases. Large uncertainties remain regarding the magnitude of the CO₂ fertilization effect [Smith *et al.*, 2016; Schimel *et al.*, 2015], but studies agree that N dynamics are a key, and often missing, component of this effect [Terrer *et al.*, 2016; Wieder *et al.*, 2015].

Climate change may affect photosynthesis and respiration in ways that are not well represented in MsTMIP models, particularly through temperature acclimation of photosynthesis and autotrophic respiration (Ra), and through carbon allocation and nutrient availability. Ra acclimates to sustained higher temperatures faster than photosynthesis [Ryan and Law, 2005; Way and Oren, 2010; Reich *et al.*, 2016], potentially increasing the fraction of photosynthesis that results in NPP. The only two MsTMIP models that allow for Ra temperature acclimation, LPJwsl and VISIT, have among the largest climate-driven increases in absolute NPP (Figure 4c). Ra acclimation was also included in the LPJml model, which was shown to reproduce the increase in ASC at ground-based stations [Sitch *et al.*, 2003; Forkel *et al.*, 2016]. Warming has also been shown to increase above-ground allocation of carbon [Poorter *et al.*, 2012; Way and Oren, 2010; Melillo *et al.*, 2011], potentially due to accelerated cycling of N through the soil that reduces the requirement for belowground carbon allocation [Melillo *et al.*, 2011; Wieder *et al.*, 2015]. Most MsTMIP models do not include N cycling, and C-N models may underestimate this effect. Aboveground carbon allocation could also drive structural changes that increase leaf area and contribute to greening, depending on the species and on ecosystem conditions [Ainsworth and Long, 2005; Way and Oren, 2010; Melillo *et al.*, 2011; McCarthy *et al.*, 2006].

Recent studies have suggested that land use change (LUC) could be responsible for a substantial portion of the ASC increase [Gray *et al.*, 2014; Zeng *et al.*, 2014]. Specifically, the transition from forest to intensive

Acknowledgments

This work was supported by the Grantham Institute: Climate Change and the Environment—Science and Solutions for a Changing Planet DTP, grant NE/L002515/1 and is a contribution to the AXA Chair Programme in Biosphere and Climate Impacts. All data used for this analysis are publicly available with full sources detailed in the supporting information. We thank Peter Rayner and an anonymous reviewer for their helpful comments on the manuscript. Funding for the Multiscale synthesis and Terrestrial Model Intercomparison Project (MSTMIP; <http://nacp.ornl.gov/MSTMIP.shtm>) activity was provided through NASA ROSES grant NNX10AG01A. Data management support for preparing, documenting, and distributing model driver and output data was performed by the Modeling and Synthesis Thematic Data Center at Oak Ridge National Laboratory (ORNL; <http://nacp.ornl.gov>), with funding through NASA ROSES grant NNX10AN681. Finalized MSTMIP data products are archived at the ORNL DAAC (<http://daac.ornl.gov>). Biome-BGC code was provided by the Numerical Terradynamic Simulation Group at University of Montana. The computational facilities were provided by NASA Earth Exchange at NASA Ames Research Center. CLM4 and GTEC simulations were supported in part by the U.S. Department of Energy (DOE), Office of Science, Biological and Environmental Research. Oak Ridge National Laboratory is managed by UT-BATTELLE for DOE under contract DE-AC05-00OR22725. CLM4-VIC research is supported in part by the U.S. Department of Energy (DOE), Office of Science, Biological and Environmental Research. PNNL is operated for the U.S. DOE by BATTELLE Memorial Institute under contract DE-AC06-76RLO1830. DLEM developed in International Center for Climate and Global Change Research, Auburn University has been supported by NASA Interdisciplinary Science Program (IDS), NASA Land Cover/Land Use Change Program (LULUC), NASA Terrestrial Ecology Program, NASA Atmospheric Composition Modeling and Analysis Program (ACMAP), NSF Dynamics of Coupled Natural-Human System Program (CNH), Decadal and Regional Climate Prediction using Earth System Models (EaSM), DOE National Institute for Climate Change Research, USDA AFRI Program, and EPA STAR program. LPJwsl work was conducted at LSCE, France, using a modified version of LPJ version 3.1 model, originally made available by the Potsdam Institute for Climate Impact Research. ORCHIDEE is a global land surface model developed

high-yielding crops could increase ASC because such crops can have more concentrated growing seasons and higher NPP than forests [Gray *et al.*, 2014]. In MSTMIP models, LUC is not a large contribution to ASC change (Figure S4), accounting for –8 to 47% of the increase in modeled ASC (Figure 4a). MSTMIP models may not be adequately representing LUC since most models do not explicitly represent crops or land management (Table S1). However, the LPJml model used by [Forkel *et al.*, 2016] does explicitly include crops and also finds a small contribution to ASC change from LUC (7%) [Forkel *et al.*, 2016]. LUC is unlikely to explain much of the large differences between models and observations in any case, particularly because most LUC has been at 30–45°N [Gray *et al.*, 2014], south of the boreal forest region that contributes the most to observed ASC changes [Graven *et al.*, 2013]. Other vegetation changes, such as the observed northward shift of the tree line [Harsch *et al.*, 2009; Elmendorf *et al.*, 2012], may have increased NPP at higher latitudes. This shift is an important driver of the increased ASC in LPJml [Forkel *et al.*, 2016] but may not be captured by MSTMIP models.

Several factors may be contributing to the observed increase in ASC over the last 50 years. We conclude that a key factor is an increase in LUE that is larger than simulated by current models. Warming and increases in atmospheric CO₂ may have caused stronger increases in LUE than in the MSTMIP models through a combination of CO₂ fertilization, respiration acclimation, increases in the rate of N cycling, and a shift to aboveground carbon allocation. Improving the modeled LUE response to CO₂ and climate change is important not only for process-based models such as those participating in MSTMIP but also for diagnostic models that use satellite-derived fAPAR to estimate NPP. This study highlights that combining atmospheric CO₂ measurements and observed greening trends provides a powerful constraint on terrestrial biosphere models, particularly through the analysis of structural and physiological ecosystem changes using the light use efficiency concept and should be included in future model benchmarking exercises.

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at the IPSL institute in France. The simulations were performed with the support of the GhG Europe FP7 grant with computing facilities provided by "LSCE" or "TGCC." SiB3 research was carried out at the Jet Propulsion Laboratory, California Institute of Technology, under a contract with the National Aeronautics and Space Administration. TEM6 research is supported in part by the U.S. Department of Energy (DOE), Office of Science, Biological and Environmental Research. Oak Ridge National Laboratory is managed by UT-BATTELLE for DOE under contract DE-AC05-00OR22725. VISIT was developed at the National Institute of Environmental Studies, Japan. This work was mostly conducted during a visiting stay at Oak Ridge National Laboratory. This study was supported by KAKENHI Grand No. 26281014 by the Japan Society of Promotion of Science.

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