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First Printing 4M, October 1980

Information contained herein is available to all persons without regard to race, color, sex, or national origin.
THERE ARE approximately 60,000 man-made ponds in Alabama which comprise an estimated 180,000 acres. These ponds are used for many purposes including sport fishing, commercial fish production, irrigation, stock watering, fire protection, water sports, and landscaping.

The chemistry of water is especially important in ponds that are managed for sport fish, bait minnow, or food fish production. Boyd (4) indicated that ponds with soft, acid waters need to be limed to enhance the effectiveness of inorganic fertilizers that are applied to increase plankton and fish production. Ponds with hard waters may not respond well to certain fertilization techniques because phosphorus applied in fertilizer is quickly precipitated as calcium phosphate (6). Ponds with low hardness and high alkalinity are often unsuitable for fertilization because of high pH resulting from over abundant phytoplankton growth and associated carbon dioxide depletion (8). In unfertilized ponds, natural productivity is often a function of alkalinity (6, 16).

Soil scientists and other agriculturists have long recognized the importance of soil surveys as an aid to crop and land management practices. However, there have been few surveys of pond waters to provide data for use in fish management. Boyd and Walley (7) found that total hardness and total alkalinity of pond waters in Alabama were closely related to the geologic and edaphic characteristics of watersheds. Waters from some physiographic regions usually had higher alkalinity and hardness values than waters from other regions. Nevertheless, considerable variation in alkalinity and hardness was encountered within a particular

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1Graduate student and Professor, respectively, Department of Fisheries and Allied Aquacultures.
region. Boyd (5) found that nitrate, total ammonia nitrogen, potassium, soluble orthophosphate, and total phosphorus concentrations in waters of Alabama ponds did not differ between physiographic regions. However, waters from ponds in pastures almost invariably had greater concentrations of nitrogen, phosphorus, and potassium than waters from ponds in woodlands.

Other than the aforementioned studies, few data are available on water chemistry of Alabama ponds. The present study was initiated to provide data on water chemistry in ponds of different soils areas in Alabama. The variables measured were total alkalinity, total hardness, calcium, magnesium, sodium, potassium, sulfate, chloride, total dissolved solids, specific conductance, iron, manganese, zinc, and copper. No attempt was made to measure pH, phosphorus, nitrate, and ammonia because the analytical values for these substances change quickly during sample storage.

MATERIALS AND METHODS

Water Samples

Water samples were collected from five of the seven soils areas of Alabama. The areas of collection were: Limestone Valleys and Uplands, Appalachian Plateau, Piedmont Plateau, Prairies, and Coastal Plains. Sampling did not include the major Flood Plains and Terraces and the Coastal Marshes because there are relatively few ponds in these soils areas. Ponds were usually located randomly from county maps provided by U.S. Soil Conservation Service workers in each county. However, in a few areas samples were simply located by sight from highways. The locations of ponds were recorded by distance and direction from landmarks (city limit signs, intersections of highways, etc.). These locations were then plotted as accurately as possible on a general soils map of Alabama (11). Ponds were not sampled unless they could definitely be assigned to a particular soils area of the soils map. The surface area of each pond was obtained from the pond owner and the type of watershed (pasture, woodland, or cropland) was recorded.

Samples were collected between March and September, 1979. Each pond was sampled one time. Water was collected by holding a 2-liter polyethylene bottle 10-cm beneath the surface and allowing it to fill. Samples were analyzed as soon after collection as possible, usually within 2 days or less.
Water Analyses

Procedures

All analytical work conformed to guidelines presented by the American Public Health Association et al. (3). Total hardness was determined by titration with 0.01M ethylenediamine tetra-acetic acid (EDTA) using Eriochrome Black T as the indicator. Calcium hardness was measured by titration to the end point of murexide with 0.01M EDTA. Total alkalinity was estimated by titration to the methyl orange end point with 0.02N sulfuric acid. Calcium concentrations were calculated from calcium hardness (6). Magnesium hardness was taken as the difference between total hardness and calcium hardness and magnesium concentrations were calculated from magnesium hardness as suggested by Boyd (6). Specific conductance was determined with a YSI Model 31 conductivity bridge.

Samples for other analyses were filtered through glass fiber filters (Gelman Type A-E) to remove suspended particles. Sulfate was determined turbidimetrically by the barium chloride method, the barium sulfate turbidity being measured with a Coleman Model 9 Nepho-colorimeter. Chloride was titrated to the diphenylcarbazone end point with standard mercuric nitrate. Total dissolved solid concentrations were estimated from the weight of the residue remaining after evaporation of 100 ml of filtrate at 105 degrees C. Sodium, potassium, iron, manganese, zinc, and copper concentrations were determined on filtrates by the Auburn University Soil Testing Laboratory using Perkin-Elmer Models 373 and 460 atomic absorption spectrophotometers. A series of 40 samples was also analyzed for calcium and magnesium by atomic absorption spectrophotometry for comparison with titrimetric data.

Quality Control

At the beginning of the study, five 12-liter water samples that covered the concentration ranges expected for each variable were each analyzed seven times by the methods listed above. These data were used to determine the precision of the procedures according to the U.S. Environmental Protection Agency (22). Afterwards, 10 percent of the samples were analyzed in duplicate. Spike-recovery tests for accuracy were not conducted. However, data for 32 randomly selected samples were subjected to two tests, anion-cation balance (12) and agreement between
calculated and measured specific conductance (14), to provide estimates of accuracy.

**RESULTS AND DISCUSSION**

Samples were collected from 203 ponds, figure 1. These included 57 from the Coastal Plains, 39 from the Prairies, 34 from the Appalachian Plateau, 36 from the Limestone Valleys and Up-

![Map of Alabama showing pond locations by soil type](image)

**FIG. 1. Location of ponds in Alabama by soils area. Map after Hajek, Gilbert, and Steers (11).**
lands, and 37 from the Piedmont Plateau. Ponds ranged in size from 0.25 to 20 hectares. Some of the ponds had received fertilizer applications, but none had been limed. The watersheds were of three types: row crops (16 ponds), pastures (98 ponds), and woodlands (89 ponds).

**Quality Control**

Precision estimates were based on analyses of five samples. The maximum standard deviations and 95 percent confidence intervals for individual measurements (6) for each variable are listed in table 1. For example, total alkalinity measurements were expected to have standard deviations of not more than ±0.24 mg/liter of means and individual measurements were expected to be within ±0.58 mg/liter of means. Duplicate analyses of samples for quality control always fell within the expected confidence intervals. Average coefficients of variation for all five samples are also presented in table 1.

Precision was relatively good for most procedures, and especially for total alkalinity, total hardness, and specific conductance, table 1. However, precision was poor for iron, manganese, zinc, and copper.

There was good agreement between calcium and magnesium concentrations determined by atomic absorption spectropho-

<table>
<thead>
<tr>
<th>Procedure</th>
<th>Maximum standard deviation</th>
<th>Maximum 95% confidence intervals for individual measurements</th>
<th>Average coefficient of variation (%)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Total alkalinity</td>
<td>0.237</td>
<td>0.58</td>
<td>0.64</td>
</tr>
<tr>
<td>Total hardness</td>
<td>0.450</td>
<td>1.10</td>
<td>0.69</td>
</tr>
<tr>
<td>Calcium</td>
<td>0.151</td>
<td>0.37</td>
<td>1.30</td>
</tr>
<tr>
<td>Magnesium</td>
<td>0.327</td>
<td>0.80</td>
<td>4.54</td>
</tr>
<tr>
<td>Potassium</td>
<td>0.181</td>
<td>0.44</td>
<td>5.53</td>
</tr>
<tr>
<td>Sodium</td>
<td>0.214</td>
<td>0.52</td>
<td>4.27</td>
</tr>
<tr>
<td>Chloride</td>
<td>0.275</td>
<td>0.67</td>
<td>3.06</td>
</tr>
<tr>
<td>Sulfate</td>
<td>0.345</td>
<td>0.84</td>
<td>5.36</td>
</tr>
<tr>
<td>Total dissolved solids</td>
<td>2.83</td>
<td>6.92</td>
<td>3.63</td>
</tr>
<tr>
<td>Specific conductance</td>
<td>2.41</td>
<td>5.90</td>
<td>0.49</td>
</tr>
<tr>
<td>Iron</td>
<td>0.048</td>
<td>0.12</td>
<td>19.3</td>
</tr>
<tr>
<td>Manganese</td>
<td>0.028</td>
<td>0.067</td>
<td>35.8</td>
</tr>
<tr>
<td>Zinc</td>
<td>0.015</td>
<td>0.036</td>
<td>19.9</td>
</tr>
<tr>
<td>Copper</td>
<td>0.009</td>
<td>0.022</td>
<td>39.1</td>
</tr>
</tbody>
</table>
tometry and by EDTA titration, figures 2 and 3. Therefore, data obtained by the titrametric technique were used throughout the study.

Good agreement was achieved between the summation in meq/liter of major anions ($\text{CO}_3^{2-}$, $\text{HCO}_3^-$, $\text{SO}_4^{2-}$, and $\text{Cl}^-$) and major cations ($\text{Ca}^{2+}$, $\text{Mg}^{2+}$, $\text{K}^+$, and $\text{Na}^+$) for randomly selected samples, figure 4. The major ions account for almost all of the total ions in natural water and the summation of anions and cations must be equal because of the principle of electrical neutrality (13, 15, 17, 21, 23). Therefore, close agreement between the summation of anions and cations in the results of a water analysis for major ions suggests a high degree of accuracy in the analytical work (12).

The sum of $\text{CO}_3^{2-}$ and $\text{HCO}_3^-$ was obtained from the total alkalinity.

$$Y = -0.76 + 1.04X$$
$$r^2 = 0.98$$

FIG. 2. Comparison between calcium concentrations determined by atomic absorption spectrophotometry and by titration with ethylenediamine tetraacetic acid. The solid line represents perfect agreement and the dashed line is the actual regression line.
The major cations and anions also account for essentially all of the electrical conductance of a water sample. The expected specific conductances for 32 randomly selected samples were calculated by the equations given by Laxen (14). These values agreed closely with measured specific conductance, figure 5, also suggesting that the analytical work was highly accurate.

**Water Chemistry of the Soils Areas**

A thorough description of the soils areas of Alabama is presented by Hajek et al. (11). Each major soils area is subdivided into different soil associations, but no attempt was made to identify the soil association or associations of the watershed of each pond. The Limestone Valleys and Uplands, Appalachian Plateau, and Prairies contain large areas of soils which developed
in material weathered from limestone. Few soils in the Coastal Plains and none in the Piedmont Plateau developed in limestone. The soils in the Piedmont Plateau developed from the oldest surface rock in Alabama. Heavy clay surface soils are especially common in the Prairies while sandy surface soils are common in the Coastal Plains.

Data for each water quality variable from ponds of all soils areas were generally skewed, so standard deviations were of no value in describing the distributions. Frequency distributions were plotted for each variable by soils area, figures 6 to 16. Even though data were skewed, standard errors of the means were computed because skewedness does not invalidate the use of the standard error (19). Means and standard errors for each variable will usually be presented in the text rather than in tables.
Total Alkalinity

Total alkalinity is a measure of the total titratable bases in water expressed as mg/liter of equivalent calcium carbonate. In natural waters, $\text{HCO}_3^-$ and $\text{CO}_3^{2-}$ are the predominate bases and total alkalinity is considered a measure of the combined concentration of these two ions (3, 6, 12, 17, 21). Total alkalinity in mg/liter as CaCO$_3$ may be converted to meq/liter of $\text{CO}_3^{2-}$ plus $\text{HCO}_3^-$ by dividing by 50 mg of CaCO$_3$ per meq.

Means and standard errors for total alkalinity values (mg/liter as CaCO$_3$) for waters of the different soils areas were: Prairies, $51.1 \pm 4.52$; Limestone Valleys and Uplands, $42.2 \pm 5.33$; Appalachian Plateau, $18.9 \pm 1.76$; Coastal Plains, $13.2 \pm 0.95$; and Piedmont Plateau, $11.6 \pm 0.79$. The greater concentrations of total alkalinity in waters of the Prairies and the Limestone Valleys and
Uplands reflect the frequent occurrence of limestone in soils of ponds and watersheds.

Few pond waters from the Prairies, figure 6, had less than 20 mg/liter total alkalinity, while a greater percentage of the pond waters from the Limestone Valleys and Uplands had less than 20 mg/liter. However, values above 40 mg/liter were common in both of these soils areas. Waters of ponds in the other three soils areas usually contained less than 20 mg/liter total alkalinity and many had values of 10 mg/liter or less. These observations on total alkalinity agree with results of Boyd and Walley (7).

Ponds with waters which contain less than 20 mg/liter total alkalinity need to be limed to ensure response to inorganic fertilization (4). Data presented in figure 6 suggest that the decision to lime a pond should be based on a water analysis for total alkalinity. For example, most waters of the Coastal Plains had less than 20 mg/liter total alkalinity, but a few had more than 20 mg/liter and would not need lime. Conversely, most pond waters of the...
Prairies had more than 20 mg/liter total alkalinity, but a few which were located on acid soils had less than 20 mg/liter and would need lime.

**Total Hardness**

Total hardness is a measure of the total concentration of divalent cations in water expressed as mg/liter as equivalent calcium carbonate. Since Ca\(^{2+}\) and Mg\(^{2+}\) are the primary divalent cations in most natural waters, total hardness is considered an estimate of the concentration of these two cations \((6, 17)\).

Means and standard errors for total hardness values (mg/liter as CaCO\(_3\)) were: Prairies, 55.5 ± 5.19; Limestone Valleys and Uplands, 49.2 ± 6.20; Appalachian Plateau, 22.0 ± 2.25; Coastal

![Histogram of Total Hardness](image)

**FIG. 7.** Frequency distribution histograms for total hardness values of waters from ponds of different soils areas in Alabama.
Plains, 12.9 ± 0.92; and Piedmont Plateau, 12.3 ± 2.02. As with total alkalinity, the high concentrations of total hardness in pond waters from the Prairies and the Limestone Valleys and Uplands reflect the large concentrations of limestone in soils. Most total hardness values were above 20 mg/liter in ponds of the Prairies, Limestone Valleys and Uplands, and Appalachian Plateau, figure 8.

**FIG. 8.** Frequency distribution histograms for calcium concentrations of waters from ponds of different soils areas in Alabama.
However, most values were below 20 mg/liter for waters of the other two soils areas. These findings agree well with those of Boyd and Walley (7). Waters with less than 20 mg/liter total hardness should be limed for effective fertilization and sportfish production (4).

**Calcium**

Means and standard errors for calcium concentrations in mg/liter were: Prairies, 19.7 ± 1.99; Limestone Valleys and Uplands, 11.9 ± 1.76; Appalachian Plateau, 5.0 ± 0.53; Coastal Plains, 3.40 ± 0.25; and Piedmont Plateau, 2.71 ± 0.19. Most values were

![Frequency distribution histograms for magnesium concentrations of waters from ponds of different soils areas in Alabama.](image-url)
above 6 mg/liter for the Prairies and Limestone Valleys and Uplands, but below 6 mg/liter for other soils areas, figure 8. High concentrations of calcium resulted from abundant limestone in soils.

**Magnesium**

Means and standard errors for magnesium concentrations in mg/liter were: Limestone Valleys and Uplands, 4.7 ± 0.91; Appalachian Plateau, 2.8 ± 0.44; Prairies, 1.5 ± 0.22; Piedmont Plateau, 1.4 ± 0.10; and Coastal Plains, 1.1 ± 0.10. Several samples from the Limestone Valleys and Uplands had more than 10 mg/liter of magnesium, figure 9. However, most waters from this soils area and from the other four areas had less than 4 mg/liter of magnesium.

**Potassium**

Means and standard errors for potassium in mg/liter were: Limestone Valleys and Uplands, 3.2 ± 0.84; Coastal Plains, 2.8 ± 0.34; Appalachian Plateau, 1.7 ± 0.37; Prairies, 1.5 ± 0.20; and Piedmont Plateau, 1.4 ± 0.12. Distributions were relatively similar and differences in means resulted primarily from extreme values, figure 10. Almost all samples contained less than 4 mg/liter potassium, but only 23 percent of the samples had less than 1 mg/liter potassium. When presented by soils areas, the percentages are: Prairies, 5; Coastal Plains, 16; Piedmont Plateau, 19; Limestone Valleys and Uplands, 33; and Appalachian Plateau, 47. Unless ponds contain more than 1 mg/liter potassium, fertilization with potassium fertilizers is probably advisable (9).

**Sodium**

Means and standard errors for sodium in mg/liter were: Prairies, 4.3 ± 0.46; Limestone Valleys and Uplands, 4.2 ± 1.81; Appalachian Plateau, 2.9 ± 0.27; Coastal Plains, 2.9 ± 0.24; and Piedmont Plateau, 2.6 ± 0.11. Distributions were similar for the five areas and most samples had less than 5 mg/liter of sodium, figure 11. Differences in means for the five areas resulted primarily from the extreme values.

**Chloride**

Means and standard errors for chloride concentrations in mg/liter were: Prairies, 6.8 ± 4.1; Limestone Valleys and Uplands, 6.6 ± 3.7; Coastal Plains, 5.5 ± 0.79; Appalachian Plateau, 3.2 ± 0.36; and Piedmont Plateau, 2.6 ± 0.29. Most samples from the five areas had less than 6 mg/liter chloride, figure 12. A few
FIG. 10. Frequency distribution histograms for potassium concentrations of waters from ponds of different soils areas in Alabama.
FIG. 11. Frequency distribution histograms for sodium concentrations of waters from ponds of different soils areas in Alabama.
samples from the Prairies and the Coastal Plains had 12 mg/liter or more of chloride. Variation between ponds was particularly great for the Prairies and the Limestone Valleys and Uplands.

**Sulfate**

Means and standard errors for sulfate concentrations in mg/liter were: Appalachian Plateau, 6.6 ± 1.04; Prairies, 4.3 ± 0.80; Limestone Valleys and Uplands, 4.2 ± 0.78; Piedmont Plateau, 1.4 ± 0.31; and Coastal Plains, 1.8 ± 0.41. A large proportion of the
FIG. 13. Frequency distribution histograms for sulfate concentrations of waters from ponds of different soils areas in Alabama.
samples from all soils areas contained less than 2 mg/liter sulfate, figure 13. With a few exceptions, waters from ponds of the Coastal Plains and Piedmont Plateau contained less than 6 mg/liter of sulfate. For the other three areas, an appreciable number of samples had sulfate concentrations of 6 to 22 mg/liter.

Iron, Manganese, Zinc, and Copper

There were no apparent differences between the mean values or distributions of these four trace elements for the five soils areas. Therefore, data for all soils areas were combined. The means and standard errors are presented in table 2 and the frequency distribution histograms in figure 14. Concentrations of iron were usually below 0.5 mg/liter and values for zinc were normally below 0.25 mg/liter. Most copper and manganese concent-

![Frequency distribution histograms for iron, manganese, copper, and zinc concentrations of water from ponds in Alabama (all soils areas combined).](image-url)
表2. 铁、锰、锌和铜在阿拉巴马州203个池塘中水中的平均浓度和标准误差（所有土壤区域合并）

<table>
<thead>
<tr>
<th>变量</th>
<th>平均浓度 (mg/liter)</th>
<th>标准误差 (mg/liter)</th>
</tr>
</thead>
<tbody>
<tr>
<td>铁</td>
<td>0.20</td>
<td>0.021</td>
</tr>
<tr>
<td>锰</td>
<td>0.03</td>
<td>0.003</td>
</tr>
<tr>
<td>锌</td>
<td>0.07</td>
<td>0.006</td>
</tr>
<tr>
<td>铜</td>
<td>0.03</td>
<td>0.005</td>
</tr>
</tbody>
</table>

浓度均低于0.05 mg/liter。这些值代表了池塘水中过滤物中痕量元素的总浓度。大量痕量元素与无机和有机化合物（18, 20, 21）络合，因此本研究中呈现的浓度远大于这些痕量元素的真实离子浓度。

**总溶解固体和电导率**

表3呈现了两种变量的平均值和标准误差，以及图15和图16的分布频率图。最高等于的溶解固体和电导率值来自普拉里亚和石灰岩谷地和高地的水，而最低值则通常来自阿巴拉契亚高地和平原。

特定电导率值对于离子浓度最大的样品最大，因为溶液的导电能力与离子浓度成正比。总溶解固体浓度也随着离子浓度的增加而增加。可溶性有机物，主要是未电荷的，也贡献于总溶解固体。因此，水中溶解有机物的含量影响总溶解固体和特定电导率的关系。

表3. 总溶解固体浓度和电导率值

<table>
<thead>
<tr>
<th>土壤区域</th>
<th>总溶解固体浓度 (mg/liter)</th>
<th>特定电导率 (μmhos/cm)</th>
</tr>
</thead>
<tbody>
<tr>
<td>地面面积</td>
<td>数量</td>
<td>平均</td>
</tr>
<tr>
<td>普拉里亚高地</td>
<td>37</td>
<td>34.5</td>
</tr>
<tr>
<td>沿海平原</td>
<td>57</td>
<td>44.3</td>
</tr>
<tr>
<td>阿巴拉契亚高地</td>
<td>34</td>
<td>60.2</td>
</tr>
<tr>
<td>石灰岩谷地和高地</td>
<td>36</td>
<td>112</td>
</tr>
<tr>
<td>普拉里亚</td>
<td>39</td>
<td>94.4</td>
</tr>
</tbody>
</table>
Waters with low alkalinity (soft, acid waters) on wooded watersheds often contain high concentrations of tannins, lignins, and other organic substances. Such waters are heavily stained (have the appearance of tea or weak coffee) and often have fairly large concentrations of total dissolved solids in comparison to specific conductance values. Waters with high concentrations of divalent ions (hard waters) seldom contain high concentrations of organic substances because divalent ions favor precipitation of colloidal organic substances and the higher pH
favors decomposition of dissolved and colloidal organic matter by bacteria (23).

**Relationships Between Variables**

Pie diagrams (12) were prepared to give a visual comparison of the ionic concentrations (meq/liter) and the proportions of major
ions in pond waters of the five soils areas, figure 17. The waters of the Prairies and the Limestone Valleys and Uplands had much greater concentrations of ions than the waters of the other areas. The waters of the Piedmont Plateau were the most dilute in ions.
Waters of the Coastal Plains and Appalachian Plateau were more similar in total ionic concentration to the waters of the Piedmont Plateau than to the waters of the other two areas.

Bicarbonate and carbonate (as calculated from total alkalinity) and calcium and magnesium comprised more than 50 percent of the total ions in waters of all five areas, figure 17. The proportions of the major ions were remarkably similar between soils areas. Therefore, pond waters from Alabama differ primarily in total concentrations of ions rather than in the proportions of different ions.

To further illustrate relationships between ions, the summations of all ions in meq/liter (X-variable) were plotted, figure 18, against the summation of $\text{HCO}_3^-$, $\text{CO}_3^{2-}$, $\text{Ca}^{2+}$, and $\text{Mg}^{2+}$ in meq/liter (Y-variable). The summation of $\text{HCO}_3^-$, $\text{CO}_3^{2-}$, $\text{Ca}^{2+}$, and $\text{Mg}^{2+}$ may actually be obtained by summing total alkalinity and total hardness in mg/liter and dividing by 50. In figure 18 the solid line represents the case where $\text{HCO}_3^-$, $\text{CO}_3^{2-}$, $\text{Ca}^{2+}$, and $\text{Mg}^{2+}$
Table 4. Regression equations and correlation coefficients for relationships between the summation of $\text{HCO}_3^-$, $\text{CO}_3^{2-}$, $\text{Ca}^{2+}$, and $\text{Mg}^{2+}$ in meq/liter (X-variable) and the summation of total major ions in meq/liter (Y-variable). Data for waters from ponds of different soils areas in Alabama

<table>
<thead>
<tr>
<th>Soils area</th>
<th>Number of samples</th>
<th>Regression equation</th>
<th>Correlation coefficient ($r$)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Piedmont Plateau</td>
<td>37</td>
<td>$Y = 0.227 + 1.037X$</td>
<td>0.95</td>
</tr>
<tr>
<td>Coastal Plains</td>
<td>57</td>
<td>$Y = 0.227 + 1.333X$</td>
<td>0.91</td>
</tr>
<tr>
<td>Appalachian Plateau and Uplands</td>
<td>34</td>
<td>$Y = 0.11 + 1.198X$</td>
<td>0.97</td>
</tr>
<tr>
<td>Limestone Valleys and Uplands</td>
<td>36</td>
<td>$Y = 0.332 + 0.146X$</td>
<td>0.86</td>
</tr>
<tr>
<td>Prairies</td>
<td>39</td>
<td>$Y = 0.345 + 1.075X$</td>
<td>0.98</td>
</tr>
</tbody>
</table>

$^1$Total alkalinity plus total hardness (in mg/liter as CaCO$_3$) divided by 50 gives summation of $\text{HCO}_3^-$, $\text{CO}_3^{2-}$, $\text{Ca}^{2+}$, and $\text{Mg}^{2+}$ in meq/liter.

are the only ions in water and the dashed line represents the situation where the summation of $\text{HCO}_3^-$, $\text{CO}_3^{2-}$, $\text{Ca}^{2+}$, and $\text{Mg}^{2+}$ is exactly one-half of the summation of the major ions. A number of data points for the Coastal Plains fell below the dashed line, indicating that the summation of $\text{HCO}_3^-$, $\text{CO}_3^{2-}$, $\text{Ca}^{2+}$, and $\text{Mg}^{2+}$ comprised less than one-half of the total major ions. However, with five exceptions, the summation of $\text{HCO}_3^-$, $\text{CO}_3^{2-}$, $\text{Ca}^{2+}$, and $\text{Mg}^{2+}$ comprised one-half or more of the total major ions in waters from ponds of the other four soils areas. Thus, by measuring total hardness and total alkalinity in waters of Alabama ponds, one may obtain a fair estimate of the total concentration of major ions in Alabama pond waters. Regression equations for calculating the summation of total ions in meq/liter from total alkalinity plus total hardness are given in table 4.

The analytical data in mg/liter were used to calculate ratios between variables, figure 19. The ratio, total alkalinity: total hardness ranged from 0.8 (Appalachian Plateau) to 1.4 (Coastal Plains). Except for the Coastal Plains, alkalinity was slightly less than total hardness. Linear correlation coefficients between total hardness (X) and total alkalinity (Y) were Piedmont ($r = 0.79; P < 0.01$), Limestone Valleys and Uplands ($r = 0.97; P < 0.01$), Coastal Plains ($r = 0.90; P < 0.01$), Prairies ($r = 0.89; P < 0.01$), and Appalachian Plateau ($r = 0.92; P < 0.01$). It is easier in the field to make total hardness than total alkalinity determinations. For practical purposes, total hardness multiplied by 1.4 would give an estimate of total alkalinity for Coastal Plains waters, while total hardness multiplied by 0.9 would give a better estimate of total alkalinity for waters of other soils areas.

The undesirable combination of high alkalinity and low hardness was seldom encountered. Only one pond had an alkalinity
value that was at least twice as great as hardness values. Boyd et al. (8) reported high alkalinity and low hardness in certain ponds which were filled by well water that had been softened by exchange of calcium in water with sodium in the geological deposits of the aquifer.

The ratio of calcium:magnesium, figure 19, was 18 for waters of the Prairies, but the ratios were 2 to 5 for other areas. This difference undoubtedly reflects a greater proportion of calcium to magnesium in Prairies soils as compared to other soils of Alabama.

Sodium concentrations were 2 to 3 times greater than potassium concentrations, figure 19. Sulfate concentrations were about twice as high as chloride concentrations in waters of the Appalachian Plateau and Limestone Valleys and Uplands, figure 19. However, in the other areas, sulfate concentrations were less than or approximately equal to chloride concentrations.
TABLE 5. CORRELATION COEFFICIENTS (r) FOR REGRESSIONS OF TOTAL HARDNESS (X-VARIABLE) AND OTHER WATER CHEMISTRY VARIABLES (Y-VARIABLES) FOR WATERS FROM PONDS OF DIFFERENT SOILS AREAS IN ALABAMA

<table>
<thead>
<tr>
<th>Soils area</th>
<th>Number of samples</th>
<th>Total alkalinity</th>
<th>Ca&lt;sup&gt;2+&lt;/sup&gt;</th>
<th>Mg&lt;sup&gt;2+&lt;/sup&gt;</th>
<th>K</th>
<th>Na</th>
<th>Cl</th>
<th>SO&lt;sub&gt;4&lt;/sub&gt;&lt;sup&gt;2-&lt;/sup&gt;</th>
<th>TDS&lt;sup&gt;2&lt;/sup&gt;</th>
<th>Specific conductance</th>
</tr>
</thead>
<tbody>
<tr>
<td>Piedmont Plateau ........</td>
<td>37</td>
<td>0.91**</td>
<td>0.73**</td>
<td>0.91**</td>
<td>0.19</td>
<td>0.35*</td>
<td>0.25</td>
<td>0.08</td>
<td>0.41**</td>
<td>0.92**</td>
</tr>
<tr>
<td>Coastal Plains ..........</td>
<td>57</td>
<td>0.90**</td>
<td>0.94**</td>
<td>0.84**</td>
<td>0.39*</td>
<td>0.23</td>
<td>0.22</td>
<td>0.09</td>
<td>0.42**</td>
<td>0.72**</td>
</tr>
<tr>
<td>Appalachian Plateau ...</td>
<td>34</td>
<td>0.92**</td>
<td>0.87**</td>
<td>0.70**</td>
<td>0.08</td>
<td>0.53**</td>
<td>0.26</td>
<td>0.47**</td>
<td>0.45**</td>
<td>0.78**</td>
</tr>
<tr>
<td>Limestone Valleys and Uplands</td>
<td>36</td>
<td>0.97**</td>
<td>0.88**</td>
<td>0.79**</td>
<td>0.05</td>
<td>0.27</td>
<td>0.23</td>
<td>0.03</td>
<td>0.57**</td>
<td>0.38**</td>
</tr>
<tr>
<td>Prairies ................</td>
<td>39</td>
<td>0.89**</td>
<td>0.96**</td>
<td>0.44**</td>
<td>0.22</td>
<td>0.08</td>
<td>0.58**</td>
<td>0.29</td>
<td>0.12</td>
<td>0.75**</td>
</tr>
</tbody>
</table>

<sup>1</sup>Significance at the 5% and the 1% probability levels indicated by one and two asterisks, respectively.
<sup>2</sup>Total dissolved solids.

TABLE 6. CORRELATION COEFFICIENTS (r) FOR REGRESSIONS OF SPECIFIC CONDUCTANCE (X-VARIABLE) AND OTHER WATER CHEMISTRY VARIABLES (Y-VARIABLES) FOR WATERS FROM PONDS OF DIFFERENT SOILS AREAS IN ALABAMA

<table>
<thead>
<tr>
<th>Soils area</th>
<th>Number of samples</th>
<th>Total alkalinity</th>
<th>Total hardness</th>
<th>Ca&lt;sup&gt;2+&lt;/sup&gt;</th>
<th>Mg&lt;sup&gt;2+&lt;/sup&gt;</th>
<th>K</th>
<th>Na</th>
<th>Cl</th>
<th>SO&lt;sub&gt;4&lt;/sub&gt;&lt;sup&gt;2-&lt;/sup&gt;</th>
<th>TDS&lt;sup&gt;2&lt;/sup&gt;</th>
</tr>
</thead>
<tbody>
<tr>
<td>Piedmont Plateau ........</td>
<td>37</td>
<td>0.77**</td>
<td>0.93**</td>
<td>0.82**</td>
<td>0.72**</td>
<td>0.32</td>
<td>0.39*</td>
<td>0.37*</td>
<td>0.03</td>
<td>0.46**</td>
</tr>
<tr>
<td>Coastal Plains ..........</td>
<td>57</td>
<td>0.81**</td>
<td>0.72**</td>
<td>0.67**</td>
<td>0.60**</td>
<td>0.63**</td>
<td>0.65**</td>
<td>0.34*</td>
<td>0.01</td>
<td>0.67**</td>
</tr>
<tr>
<td>Appalachian Plateau ...</td>
<td>34</td>
<td>0.82**</td>
<td>0.75**</td>
<td>0.53**</td>
<td>0.37**</td>
<td>0.38*</td>
<td>0.56**</td>
<td>0.52</td>
<td>0.39*</td>
<td>0.41*</td>
</tr>
<tr>
<td>Limestone Valleys and Uplands</td>
<td>36</td>
<td>0.32</td>
<td>0.38*</td>
<td>0.42**</td>
<td>0.14</td>
<td>0.51**</td>
<td>0.46**</td>
<td>0.45**</td>
<td>0.58**</td>
<td>0.62**</td>
</tr>
<tr>
<td>Prairies ................</td>
<td>39</td>
<td>0.53**</td>
<td>0.75**</td>
<td>0.72**</td>
<td>0.37*</td>
<td>0.09</td>
<td>0.38**</td>
<td>0.97**</td>
<td>0.02</td>
<td>0.89**</td>
</tr>
</tbody>
</table>

<sup>1</sup>Significance at the 5% and 1% probability levels indicated by one and two asterisks, respectively.
<sup>2</sup>Total dissolved solids.
Since total hardness and specific conductance are both easily measured under field conditions, regression analyses were conducted to ascertain if concentrations of other ions could be estimated from either of these determinations. The regression coefficients \( r \) are summarized in tables 5 and 6.

The majority of the regression coefficients were statistically significant. This was expected since specific conductance is related to total ionic concentration, and because total hardness is closely related to total alkalinity and the ions of hardness and alkalinity usually comprised more than 50 percent of the ions in Alabama pond waters. However, a regression coefficient of 0.87 is required if a regression equation is to account for at least 75 percent of the variation in the dependent variable, so many of the regression equations are of little predictive value. The regression equations for regression analyses with \( r > 0.86 \) are tabulated in tables 7 and 8.

### Table 7. Regression Equations for Relationships Between Total Hardness (X-variable) and Selected Water Chemistry Variables (Y-variables) for Waters From Ponds of Different Soils Areas in Alabama.

<table>
<thead>
<tr>
<th>Soils area</th>
<th>Y-variable</th>
<th>Equation</th>
</tr>
</thead>
<tbody>
<tr>
<td>Piedmont Plateau</td>
<td>Total alkalinity</td>
<td>( Y = 1.440 + 0.827 \times X )</td>
</tr>
<tr>
<td></td>
<td>Magnesium</td>
<td>( Y = 0.121 + 0.102 \times X )</td>
</tr>
<tr>
<td></td>
<td>Specific conductance</td>
<td>( Y = 5.736 + 2.797 \times X )</td>
</tr>
<tr>
<td>Coastal Plains</td>
<td>Total alkalinity</td>
<td>( Y = 1.073 + 0.935 \times X )</td>
</tr>
<tr>
<td></td>
<td>Calcium</td>
<td>( Y = -0.038 + 0.242 \times X )</td>
</tr>
<tr>
<td>Appalachian Plateau</td>
<td>Total alkalinity</td>
<td>( Y = 1.473 + 0.722 \times X )</td>
</tr>
<tr>
<td></td>
<td>Calcium</td>
<td>( Y = -1.009 + 0.236 \times X )</td>
</tr>
<tr>
<td>Limestone Valleys and Uplands</td>
<td>Total alkalinity</td>
<td>( Y = 1.294 + 0.832 \times X )</td>
</tr>
<tr>
<td></td>
<td>Calcium</td>
<td>( Y = 0.223 + 0.238 \times X )</td>
</tr>
<tr>
<td>Prairies</td>
<td>Total alkalinity</td>
<td>( Y = 8.169 + 0.772 \times X )</td>
</tr>
<tr>
<td></td>
<td>Calcium</td>
<td>( Y = -2.520 + 0.401 \times X )</td>
</tr>
</tbody>
</table>

### Table 8. Regression Equations for Relationships Between Specific Conductance (X-variable) and Selected Water Chemistry Variables (Y-variables) for Waters From Ponds of Different Soils Areas in Alabama.

<table>
<thead>
<tr>
<th>Soils area</th>
<th>Y-variable</th>
<th>Equation</th>
</tr>
</thead>
<tbody>
<tr>
<td>Piedmont Plateau</td>
<td>Total hardness</td>
<td>( Y = 0.012 + 0.306 \times X )</td>
</tr>
<tr>
<td>Coastal Plains</td>
<td>None</td>
<td></td>
</tr>
<tr>
<td>Appalachian Plateau</td>
<td>None</td>
<td></td>
</tr>
<tr>
<td>Limestone Valleys and Uplands</td>
<td>None</td>
<td></td>
</tr>
<tr>
<td>Prairies</td>
<td>Chloride</td>
<td>( Y = -27.08 + 0.259 \times X )</td>
</tr>
<tr>
<td></td>
<td>Total dissolved solids</td>
<td>( Y = 17.42 + 0.479 \times X )</td>
</tr>
</tbody>
</table>
Effect of Watershed Cover

Analytical data were averaged across all soils areas for the three types of watersheds: row crops, pasture, and woods, figure 20. Although differences were noted for all variables, only the differences in potassium have practical significance. Results in figure 20 show that potassium concentrations were roughly twice as great in waters of ponds on row crop or pasture watersheds as in waters from ponds on wooded watersheds. Thus, the need for potassium fertilization is less for ponds on row crop or pasture watersheds.

Four basic types of water were noted in Alabama ponds as follows: clear water (rather unusual except in certain areas with abundant limestone), waters turbid with suspended soil particles...
(ponds with row crops on watersheds), waters turbid with plankton (fertilized ponds or ponds in well-managed pastures), and waters stained with organic substances (ponds with wooded watersheds or pasture ponds with soft, acid waters). Obviously, there was some mingling of the types. For example, fertilized ponds with heavily stained waters often contained plankton blooms.

**Ionic Concentrations and Activities**

Calculations involving the methods of thermodynamics (equilibrium calculations) require that measured ionic concentrations be reduced to ionic activities. The simplest technique for converting ionic concentrations to ionic activities for equilibrium calculations is the use of the Debye-Huckel equation (1, 10) to calculate the activity coefficients of the ions of interest. The activity coefficients are then multiplied by the measured molar concentrations to give activities. The Debye-Huckel equation is

$$-\log \gamma_i = \frac{(A) (Z_i)^2 (\mu)^{1/2}}{1 + (B) (a_i) (\mu)^{1/2}}$$

where \(A\) and \(B\) are temperature dependent constants, \(Z_i\) = ionic charge, \(a_i\) = the effective ion size, \(\mu\) = ionic strength, and \(\gamma_i\) = the activity coefficient of the ion. Values for \(A\), \(B\), and \(a_i\) are given by Garrels and Christ (10) and ionic strength can be calculated as

$$\mu = \sum \frac{(m_i) (Z_i)^2}{2}$$

where \(m_i\) = the measured molar concentration of the given ion.

A fraction of the anions and cations in aqueous solutions are strongly attracted to each other and behave as un-ionized species called "ion-pairs." Adams (11) demonstrated that ion-pairing affects ionic strength, ionic concentration, and activity coefficients. The effect of ion-pairing on the calculated activity increases with increasing concentrations of ions and is appreciable in soil solutions (2). The method for correcting for ion-pairing involves (a) measuring the ionic concentrations and calculating ionic strength assuming no ion-pairing, (b) calculating ionic activities, (c) calculating ion-pair concentrations, (d) revising ionic concentrations and ionic strength based on the calculated ion-pair values, and (e) repeating steps (b), (c), and (d) until all ionic concentrations remain unchanged with succeeding calculations. This iterative process is slow unless it is programmed into a computer.
### Table 9. Comparison of Activities (meq/liter) Calculated by the Debye-Hückel Equation Without (Initial) and With (Corrected) Correction for Ion-Pairing

<table>
<thead>
<tr>
<th>Soils area</th>
<th>HCO$_3^-$</th>
<th>Ca$^{2+}$</th>
<th>Mg$^{2+}$</th>
<th>Na$^+$</th>
<th>K$^+$</th>
<th>Cl$^-$</th>
<th>SO$_4^{2-}$</th>
</tr>
</thead>
<tbody>
<tr>
<td><strong>Piedmont Plateau</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Initial</td>
<td>0.23</td>
<td>0.061</td>
<td>0.054</td>
<td>0.11</td>
<td>0.035</td>
<td>0.071</td>
<td>0.013</td>
</tr>
<tr>
<td>Corrected</td>
<td>0.23</td>
<td>0.061</td>
<td>0.054</td>
<td>0.11</td>
<td>0.035</td>
<td>0.071</td>
<td>0.013</td>
</tr>
<tr>
<td><strong>Coastal Plains</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Initial</td>
<td>0.26</td>
<td>0.076</td>
<td>0.040</td>
<td>0.12</td>
<td>0.070</td>
<td>0.15</td>
<td>0.016</td>
</tr>
<tr>
<td>Corrected</td>
<td>0.26</td>
<td>0.076</td>
<td>0.040</td>
<td>0.12</td>
<td>0.070</td>
<td>0.15</td>
<td>0.016</td>
</tr>
<tr>
<td><strong>Appalachian Plateau</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Initial</td>
<td>0.36</td>
<td>0.11</td>
<td>0.10</td>
<td>0.12</td>
<td>0.071</td>
<td>0.087</td>
<td>0.058</td>
</tr>
<tr>
<td>Corrected</td>
<td>0.36</td>
<td>0.11</td>
<td>0.10</td>
<td>0.12</td>
<td>0.071</td>
<td>0.087</td>
<td>0.058</td>
</tr>
<tr>
<td><strong>Limestone Valleys and Uplands</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Initial</td>
<td>0.81</td>
<td>0.25</td>
<td>0.16</td>
<td>0.18</td>
<td>0.078</td>
<td>0.18</td>
<td>0.036</td>
</tr>
<tr>
<td>Corrected</td>
<td>0.80</td>
<td>0.24</td>
<td>0.16</td>
<td>0.18</td>
<td>0.078</td>
<td>0.18</td>
<td>0.034</td>
</tr>
<tr>
<td><strong>Prairies</strong></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>Initial</td>
<td>0.97</td>
<td>0.40</td>
<td>0.051</td>
<td>0.18</td>
<td>0.066</td>
<td>0.12</td>
<td>0.036</td>
</tr>
<tr>
<td>Corrected</td>
<td>0.96</td>
<td>0.39</td>
<td>0.051</td>
<td>0.18</td>
<td>0.066</td>
<td>0.12</td>
<td>0.034</td>
</tr>
</tbody>
</table>

The influence of ion-pairing on calculated activities of major ions was determined for hypothetical samples assigned average concentrations of major ions for each of the five soils areas. Activities calculated from initial data (assuming no ion-pairing) and by the iterative procedure of Adams (1) were the same for the Piedmont Plateau, Coastal Plains, and Appalachian Plateau, table 9. There were only slight differences for the two methods of calculating activities for the Limestone Valleys and Uplands and Prairies, table 9. Therefore, the Debye-Hückel procedure may be used without correcting data for ion-pairing to convert analytical values for waters from Alabama ponds to activities. It is important to convert analytical concentrations to activities since, for the hypothetical samples, activity coefficients were as low as 0.819 for divalent ions and 0.952 for monovalent ions.

**CONCLUSIONS**

The average concentrations of total ions in waters of Alabama ponds differed with soils areas as follows: Piedmont Plateau < Coastal Plains < Appalachian Plateau < Limestone Valleys and Uplands < Prairies. Waters were essentially dilute solutions of alkaline earth carbonates and bicarbonates with low concentrations of sodium, potassium, chloride, and sulfate and with minute amounts of iron, manganese, zinc, and copper. Concentrations of nitrogen and phosphorus were not measured. The major differences in ionic concentrations in waters from the different areas
resulted from variable amounts of alkaline earth carbonates and bicarbonates (total hardness and total alkalinity). Soils of areas in which ponds had high alkalinity and hardness had developed in limestone. Total ionic concentrations in waters of Alabama ponds can be estimated accurately by regression equations using the sum of total hardness and total alkalinity as the independent variable (X-variable).

For a given soils area, waters differed considerably in concentrations of all variables. Values were highly skewed, so data were presented in frequency distribution histograms. Means and standard errors for each variable and soils area were presented, but these values have limited practical value because of the great variation among individual samples.

Concentrations of total dissolved solids and values for specific conductance were often correlated with concentrations of individual ions. However, regression equations had little predictive value.

Management practices on watersheds did not greatly affect concentrations of major ions. However, ponds on wooded watersheds often had organically stained waters, ponds with row crops on their watersheds were often turbid with suspended soil particles, and pasture ponds frequently had plankton blooms.

The effect of ion-pairing on the calculation of activities from analytical data was not appreciable for waters from Alabama ponds.

LITERATURE CITED


Alabama's Agricultural Experiment Station System
AUBURN UNIVERSITY

With an agricultural research unit in every major soil area, Auburn University serves the needs of field crop, livestock, forestry, and horticultural producers in each region in Alabama. Every citizen of the State has a stake in this research program, since any advantage from new and more economical ways of producing and handling farm products directly benefits the consuming public.

Research Unit Identification

★ Main Agricultural Experiment Station, Auburn.
★ E. V. Smith Research Center, Shorter.

1. Tennessee Valley Substation, Belle Mina.
2. Sand Mountain Substation, Crossville.
4. Upper Coastal Plain Substation, Winfield.
5. Forestry Unit, Fayette County.
6. Foundation Seed Stocks Farm, Thorsby.
7. Chilton Area Horticulture Substation, Clanton.
8. Forestry Unit, Coosa County.
9. Piedmont Substation, Camp Hill.
10. Plant Breeding Unit, Tallassee.
11. Forestry Unit, Autauga County.
12. Prattville Experiment Field, Prattville.
13. Black Belt Substation, Marion Junction.
14. The Turnipseed-Ikenberry Place, Union Springs.
15. Lower Coastal Plain Substation, Camden.
16. Forestry Unit, Barbour County.
17. Monroeville Experiment Field, Monroeville.
18. Wiregrass Substation, Headland.
20. Solon Dixon Forestry Education Center, Covington and Escambia counties.
21. Ornamental Horticulture Field Station, Spring Hill.
22. Gulf Coast Substation, Fairhope.